

# Methane observations from the Greenhouse Gases Observing SATellite: Comparison to ground-based TCCON data and model calculations

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[1] We report new short-wave infrared (SWIR) column retrievals of atmospheric methane ( $X_{CH_4}$ ) from the Japanese Greenhouse Gases Observing SATellite (GOSAT) and compare observed spatial and temporal variations with correlative ground-based measurements from the Total Carbon Column Observing Network (TCCON) and with the global 3-D GEOS-Chem chemistry transport model. GOSAT  $X_{CH_4}$  retrievals are compared with daily TCCON observations at six sites between April 2009 and July 2010 (Bialystok, Park Falls, Lamont, Orleans, Darwin and Wollongong). GOSAT reproduces the site-dependent seasonal cycles as observed by TCCON with correlations typically between 0.5 and 0.7 with an estimated single-sounding precision between 0.4–0.8%. We find a latitudinal-dependent difference between the  $X_{CH_4}$  retrievals from GOSAT and TCCON which ranges from 17.9 ppb at the most northerly site (Bialystok) to –14.6 ppb at the site with the lowest latitude (Darwin). We estimate that the mean smoothing error difference included in the GOSAT to TCCON comparisons can account for 15.7 to 17.4 ppb for the northerly sites and for 1.1 ppb at the lowest latitude site. The GOSAT  $X_{CH_4}$  retrievals agree well with the GEOS-Chem model on annual (August 2009 – July 2010) and monthly timescales, capturing over 80% of the zonal variability. Differences between model and observed  $X_{CH_4}$  are found over key source regions such as Southeast Asia and central Africa which will be further investigated using a formal inverse model analysis. **Citation:** Parker, R., et al. (2011), Methane observations from the Greenhouse Gases Observing SATellite: Comparison to ground-based TCCON data and model calculations, *Geophys. Res. Lett.*, 38, L15807, doi:10.1029/2011GL047871.

## 1. Introduction

[2] Methane is the second most important anthropogenic greenhouse gas, with a radiative forcing that is comparable to  $CO_2$  over a 20-year horizon. Methane also influences tropospheric ozone and water vapor, further increasing its

importance to the Earth's radiative budget. Recent unexpected changes in surface concentrations of  $CH_4$  emphasize gaps in our current understanding of the  $CH_4$  budget [Bousquet *et al.*, 2006; Rigby *et al.*, 2008; Dlugokencky *et al.*, 2009], which has relied on highly accurate but sparse ground-based measurements. Satellite observations of  $CH_4$  offer new insights into the magnitude of regional sources and sinks and can help overcome large uncertainties associated with the upscaling and interpretation of surface concentration data.

[3] The usefulness of short-wave infrared (SWIR) measurements for inverse modeling has been demonstrated with the SCIAMACHY instrument onboard ENVISAT [Bergamaschi *et al.*, 2009; Bloom *et al.*, 2010] for which there is already a 7 year (2003–2009) time-series of global methane total column observations available with near-surface sensitivity [Frankenberg *et al.*, 2011; Schneising *et al.*, 2010].

[4] We report SWIR measurements of  $CH_4$  from the Japanese Greenhouse Gases Observing SATellite (GOSAT) (section 2), the first dedicated greenhouse gas sensor, which can extend and improve the time-series obtained from SCIAMACHY. In section 3, we describe the optimal estimation retrieval process used to retrieve  $X_{CH_4}$ . We compare these retrievals against correlative accurate/precise ground-based  $X_{CH_4}$  data (section 4), and model data (section 5). We conclude the paper in section 6.

## 2. The Greenhouse Gases Observing SATellite (GOSAT)

[5] GOSAT was launched on 23 January 2009 [Kuze *et al.*, 2009] by the Japanese Space Agency and provides global measurements of total column  $CO_2$  and  $CH_4$  from its SWIR bands with global coverage every 3 days. It is equipped with two instruments: 1) the Thermal And Near Infrared Sensor for carbon Observations – Fourier Transform Spectrometer (TANSO-FTS) and 2) the Cloud and Aerosol Imager (TANSO-CAI).

[6] The TANSO-FTS instrument has four spectral bands with a high spectral resolution ( $0.3\text{ cm}^{-1}$ ), three of which operate in the SWIR at around 0.76, 1.6 and  $2.0\text{ }\mu\text{m}$  providing sensitivity to the near-surface absorbers with the fourth channel operating in the thermal infrared between 5.5 and  $14.3\text{ }\mu\text{m}$  providing mid-tropospheric sensitivity.

[7] The measurement strategy of TANSO-FTS is optimized for the characterization of continental-scale sources and sinks. These measurements nominally consist of 5 across-track points, separated by  $\sim 100\text{ km}$ , with a ground footprint diameter of  $\sim 10.5\text{ km}$  and a 4 second exposure

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duration. Recent results from GOSAT include the preliminary validation of their CO<sub>2</sub> and CH<sub>4</sub> products against TCCON ground-based FTS instruments [Morino *et al.*, 2011].

### 3. GOSAT XCH<sub>4</sub> Retrieval Approach

[8] The OCO ‘full physics’ retrieval algorithm was developed for the NASA Orbiting Carbon Observatory (OCO) mission to retrieve  $X_{CO_2}$  (dry-air column-averaged mole fraction of CO<sub>2</sub>) from a simultaneous fit of SWIR O<sub>2</sub> and CO<sub>2</sub> bands. The OCO algorithm has been modified to operate on GOSAT spectra and we use it to perform global retrievals of  $X_{CH_4}$ .

[9] The retrieval algorithm uses an iterative retrieval scheme based on Bayesian optimal estimation to retrieve a set of atmospheric/surface/instrument parameters, referred to as the state vector, from measured spectral radiances [Boesch *et al.*, 2011; Connor *et al.*, 2008]. The forward model, used to relate the state vector to the measured radiances, includes the LIDORT radiative transfer model combined with a fast 2-orders-of-scattering vector radiative transfer code [Natraj *et al.*, 2008]. In addition, we use the low-streams interpolation functionality of the code [O’Dell, 2010] to accelerate the radiative transfer component of the retrieval algorithm.

[10] For our GOSAT CH<sub>4</sub> retrieval, we have adopted the CO<sub>2</sub> proxy method [Frankenberg *et al.*, 2011]. CO<sub>2</sub> is known to vary in the atmosphere much less than CH<sub>4</sub> and as the CO<sub>2</sub> absorption band is spectrally close to that of CH<sub>4</sub> we can use the CO<sub>2</sub> as a proxy for the light path to minimize common spectral artefacts due to aerosol scattering and instrumental effects [Butz *et al.*, 2010]. CH<sub>4</sub> and CO<sub>2</sub> retrievals are carried out sequentially with channels at 1.65  $\mu$ m and 1.61  $\mu$ m respectively. In order to obtain a volume mixing ratio (VMR) of CH<sub>4</sub>, it is necessary to multiply the  $X_{CH_4}/X_{CO_2}$  ratio by a model  $X_{CO_2}$ . We obtain the CO<sub>2</sub> VMRs from the CarbonTracker model [Peters *et al.*, 2007], convolved with scene-dependent instrument averaging kernels obtained from the GOSAT CO<sub>2</sub> retrieval. The model data is normalized such that the annual global mean is consistent with our retrieved CO<sub>2</sub>.

[11] The state vector of our retrieval consists of a 20-level profile for CH<sub>4</sub> and CO<sub>2</sub> VMRs, scaling factors for H<sub>2</sub>O VMR and temperature profiles, surface albedo and a spectral shift/stretch. Details for the a priori can be found in the auxiliary material.<sup>1</sup> We process the latest versions of the GOSAT Level 1B files (050050C to 100100C) acquired via the GOSAT User Interface Gateway and apply the recommended radiometric calibration with the noise calculated from the standard deviation of the out-of-band signal. Spectra over ocean or with a signal-to-noise ratio below 50 are excluded.

[12] The spectroscopic parameters used for CO<sub>2</sub> and O<sub>2</sub> are taken from v3.2 of the OCO linelists which include line-mixing, whilst CH<sub>4</sub> and H<sub>2</sub>O are taken from the TCCON linelists based on HITRAN08 [Rothman *et al.*, 2009] with updates to CO<sub>2</sub> [Toth *et al.*, 2008], H<sub>2</sub>O [Toth, 2005; Jenouvrier *et al.*, 2007] and CH<sub>4</sub> [Frankenberg *et al.*, 2008].

[13] Clouds shield parts of the atmosphere and consequently compromise the interpretation of retrieved  $X_{CH_4}$ . To remove cloudy scenes, we first retrieve a clear-sky surface

pressure from the O<sub>2</sub> A-Band which we compare to the ECMWF surface pressure interpolated to the measurement time and location. If they differ by more than 20 hPa we consider the scene to be cloudy and no  $X_{CH_4}$  retrieval is performed.

[14] We examined a total of 1,101,442 GOSAT measurements performed over land between August 2009 and July 2010, which after filtering for signal-to-noise, cloud and data quality (currently  $\chi^2$  only) left 197,090  $X_{CH_4}$  retrievals (see auxiliary material for details).

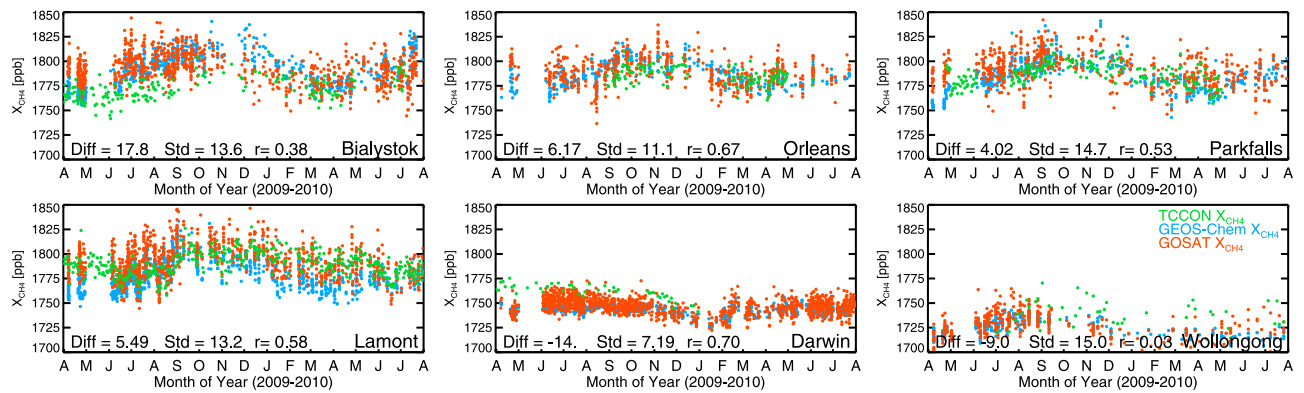
### 4. Comparison to the Total Carbon Column Observing Network (TCCON)

[15] TCCON is a global network of ground-based high resolution Fourier Transform Spectrometers recording direct solar spectra in the near-infrared spectral region [Wunch *et al.*, 2011]. The overall objectives include improving the understanding of the carbon cycle and validating satellite retrievals by providing precise and accurate observations of  $X_{CO_2}$  and  $X_{CH_4}$  for a number of locations.

[16] Figure 1 shows all GOSAT retrievals along with the daily mean TCCON measurements of  $X_{CH_4}$  for the period April 2009 to July 2010 over TCCON sites at Park Falls (USA, 45.95°N, 90.27°W), Lamont (USA, 36.60°N, 97.49°W), Darwin (Australia, 12.42°S, 130.89°E), Wollongong (Australia, 34.41°S, 150.88°E), Orleans (France, 47.97°N, 2.11°E) and Bialystok (Poland, 53.23°N, 23.03°E). These sites have been chosen with respect to their location, avoiding islands and hilly terrain and to obtain good latitudinal coverage (see map in Figure 3). The TCCON data uses GFIT Version 4.3.1 with GSETUP Version 2.7.4 and a CH<sub>4</sub> a priori profile with no geographical or temporal variation. TCCON columns are divided by a single, global factor of  $0.978 \pm 0.002$  [Wunch *et al.*, 2011], based on calibration against aircraft overflights, to compensate for some systematic spectroscopic biases. No such factor is currently applied to the GOSAT retrievals, hence we may expect a systematic bias to TCCON measurements. In addition, the data for Wollongong and Darwin were produced using updated spectroscopic parameters which may lead to a 0.15% reduction in the  $X_{CH_4}$ . Gaps in the TCCON data time-series are due to cloud and instrumental issues. In nominal mode, individual GOSAT measurements are spaced by ~100 km so that only a few GOSAT measurements are obtained directly in the proximity of a TCCON site. This situation is improved by GOSAT carrying out specific point observations over TCCON sites which increase the number of soundings over a small area. To compensate for the small number of GOSAT soundings directly over TCCON sites we include all cloud-free GOSAT measurements over land within  $\pm 5^\circ$  of each TCCON site.

[17] To perform a quantitative comparison, if more than one GOSAT measurement over a TCCON site is available for a given day, we take the mean. Similarly, the mean of TCCON measurements within a 2 hour coincidence of these GOSAT measurements is taken, providing a single GOSAT and TCCON measurement per day. The difference and correlation coefficient given in Figure 1 are inferred from the above coincident daily-averaged measurements. The single-sounding precision is estimated from the standard deviation of the individual GOSAT-TCCON differences and we obtain values ranging from 7 to 15 ppb (0.4 to 0.8%) for the different sites.

<sup>1</sup>Auxiliary materials are available in the HTML. doi:10.1029/2011GL047871.



**Figure 1.** Time-series plots showing the GOSAT and GEOS-Chem  $X_{CH_4}$  along with the daily mean TCCON  $X_{CH_4}$  for the six TCCON sites; Bialystok, Orleans, Parkfalls, Lamont, Darwin and Wollongong. The GEOS-Chem data displayed has been convolved with the GOSAT scene-specific averaging kernels and thus is only intended for comparison to the GOSAT data. The difference (ppb) and correlation coefficient are inferred only for days where a 2 hour coincidence exists between TCCON and GOSAT observations. The single-sounding precision (ppb) is estimated by calculating the standard deviation of the individual GOSAT-TCCON differences.

[18] GOSAT observations generally reproduce the seasonal cycle observed by TCCON data at each site, with correlations calculated from daily data typically between 0.5 and 0.7. The two exceptions are Bialystok ( $r = 0.4$ ) and Wollongong (where the correlation is insignificant). The amplitude of the seasonal cycle at Wollongong is small and the data are relatively sparse. Additionally, there are significant local methane sources (such as coal mining) as well as significant pollutant transport from Sydney such that a low correlation with GOSAT is not unexpected considering our spatial coincidence criteria. Additional scatter in the GOSAT retrieval could be due to local CO<sub>2</sub> sources which can impact the proxy retrieval.

[19] We find a latitudinal-dependent difference between the GOSAT  $X_{CH_4}$  retrievals and the TCCON data which ranges from 17.9 ppb at the most northerly site (Bialystok) to -14.6 ppb at the site with the lowest latitude (Darwin).

[20] The ground based TCCON retrieval and the space-based GOSAT retrieval have been carried out using different a priori methane profiles and their vertical sensitivity as described by their averaging kernels will differ. Consequently the observed differences between the GOSAT and TCCON retrievals will include a contribution from smoothing error differences [Rodgers and Connor, 2003]. We assess the effect of the smoothing errors for GOSAT and TCCON retrievals by applying the averaging kernels for both retrievals to methane profiles from the GEOS-Chem model sampled for each sounding. The inferred mean smoothing error difference is largest for Bialystok and Orleans with values of 15.7 and 17.4 ppb respectively and smallest for Darwin with a value of 1.1 ppb. Taking into account this effect, we estimate that the biases in our GOSAT retrieval range from -17 to 2 ppb. A table with the results of the comparison is given in the auxiliary material.

## 5. Comparison to the GEOS-Chem Chemistry Transport Model

[21] We use the GEOS-Chem global 3-D chemistry transport model (v8-01-01), driven by GEOS-5 assimilated meteorology from the Global Modeling and Assimilation

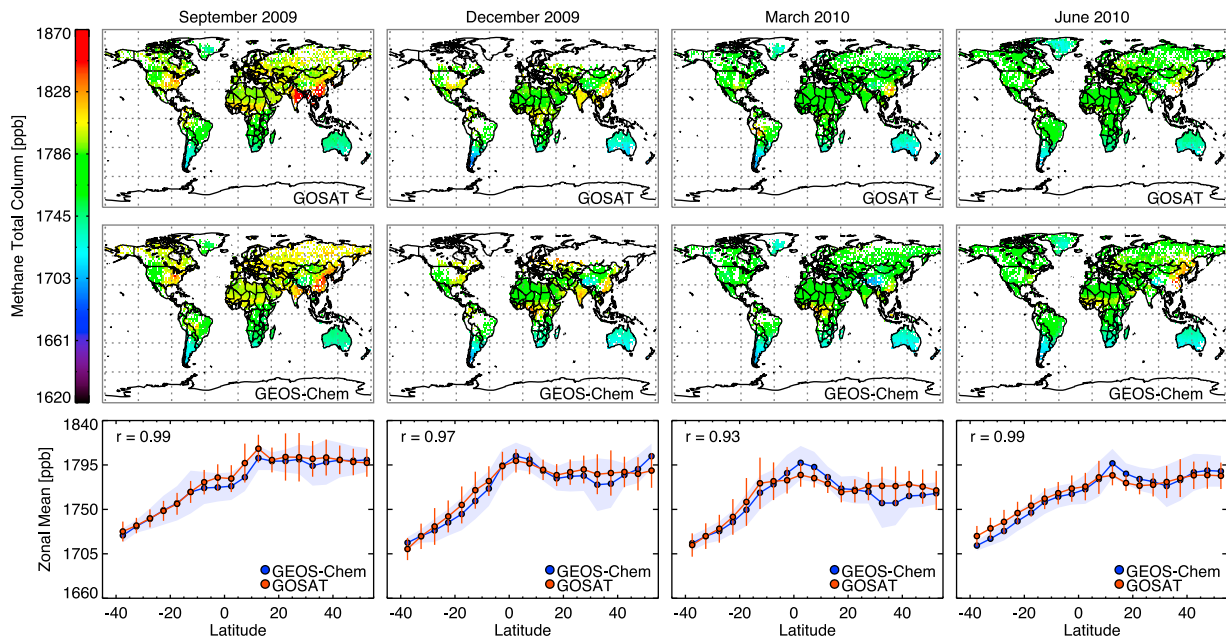
Office based at NASA Goddard, to relate CH<sub>4</sub> surface sources and sinks to atmospheric concentrations (A. Fraser et al., The Australian methane budget: Interpreting surface and train-borne measurements using a chemistry transport model, submitted to *Journal of Geophysical Research*, 2011). The 3-D meteorological data are updated every 6 hours and mixing depths and surface fields are updated every 3 hours. We use these data at a horizontal resolution of 2° latitude × 2.5° longitude, and with 47 hybrid vertical level ranging from the surface to the mesosphere, 30 of which are below 12 km.

[22] In the absence of more recent information, we use EDGAR and biomass burning from 2008 for 2009–2010; and use rice/wetland emissions from 2007 for 2008–2010 (see auxiliary material for details). Given the large uncertainties associated with any individual source estimates of CH<sub>4</sub> we anticipate that this assumption will only affect the magnitude and not the location of the emissions. The model is sampled at the time and location of individual GOSAT measurements, and is convolved with the GOSAT measurement specific averaging kernels.

[23] Figure 2 presents model and GOSAT monthly global distributions of  $X_{CH_4}$  for September and December of 2009 and March and June of 2010; other monthly mean comparisons can be found in the auxiliary material. The resulting model and observed values range from 1680 ppb to 1810 ppb, with elevated values over known continental source regions of CH<sub>4</sub>.

[24] There is very good ( $0.93 \leq r \leq 0.99$ ) agreement between model and observed zonal mean hemispheric gradients, varying month to month, as expected with seasonal sources and sinks, with the model reproducing most of the observed variability.

[25] Differences between model and observed  $X_{CH_4}$  values of up to 20 ppb are found over latitude bands which include strong emission sources. Figure 2 also shows model and GOSAT monthly spatial distributions. Large seasonal values over India and parts of southwest China, peaking in September, are due mainly to rice paddy emissions. We find large values over parts of Alaska, Canada and boreal Asia during northern hemisphere summer months, indicative of wetland



**Figure 2.** Monthly mean gridded  $2^{\circ} \times 2^{\circ}$  maps of the (top) GOSAT  $X_{CH_4}$  and (middle) GEOS-Chem  $X_{CH_4}$  for September 2009, December 2009, March 2010 and June 2010 sampled at the GOSAT measurement times and locations with the GOSAT scene-specific averaging kernel applied. (bottom) The zonal means between  $40^{\circ}\text{S}$  and  $50^{\circ}\text{N}$  in  $5^{\circ}$  bins for these months are also shown with the error bars (red) and shaded area (blue) represent the standard deviation of the GOSAT and GEOS-Chem data respectively. Also given is the correlation coefficient.

emissions and to a lesser extent wildfires. We also find elevated columns over known biomass burning regions (e.g., Africa, South America). As the relative enhancement in CH<sub>4</sub> and CO<sub>2</sub> due to biomass burning is similar, the proxy retrieval method may have reduced sensitivity to biomass burning emissions under certain conditions. Additionally, biomass burning emissions of CH<sub>4</sub> will be accompanied by combustion-related aerosol which may compromise the quality of retrieved columns and are removed by the cloud screening, helping to explain the relative sparseness of coverage over such regions.

[26] Figure 3 shows GOSAT and model  $X_{CH_4}$  time-series, for the globe and for individual geographic regions of interest: southern Africa, northern Africa, Australia, the Amazon basin, North America, Southeast Asia, and Russia. The model and GOSAT global seasonal cycles, peaking in September and reaching a minimum in March with an amplitude of 25 ppb, agree to within a fraction of a percent with the model reproducing >92% of the observed variability.

[27] We have chosen contrasting geographical regions that reflect different seasonal sources. The model generally captures more than 70% of the observed variability, but can only capture half of the observed variability over the Amazon basin, which we attribute to known weaknesses in wetland emission models [e.g., Bloom *et al.*, 2010].

[28] The northern African region incorporates only minor sources and therefore provides our control, representing changes only in the seasonal CH<sub>4</sub> background due to the OH sink. For this region, we find a very high agreement ( $r = 0.97$ ) in the magnitude and phase of  $X_{CH_4}$  between the measurement and model. Australia, similarly, is relatively isolated from large-scale CH<sub>4</sub> sources and again we find consistency between model and measurement ( $r = 0.85$ ) in capturing the small variability (amplitude of 20 ppb).

[29] North America, whilst including a range of sources, shows good agreement but with the amplitude of the seasonal cycle slightly higher (by  $\sim 10$  ppb) in the model. The southern Africa region is dominated by wetland emissions and biomass burning. The model reproduces almost 70% of the variability observed by GOSAT but values tend to be smaller by  $\sim 10$  ppb for most of the year.

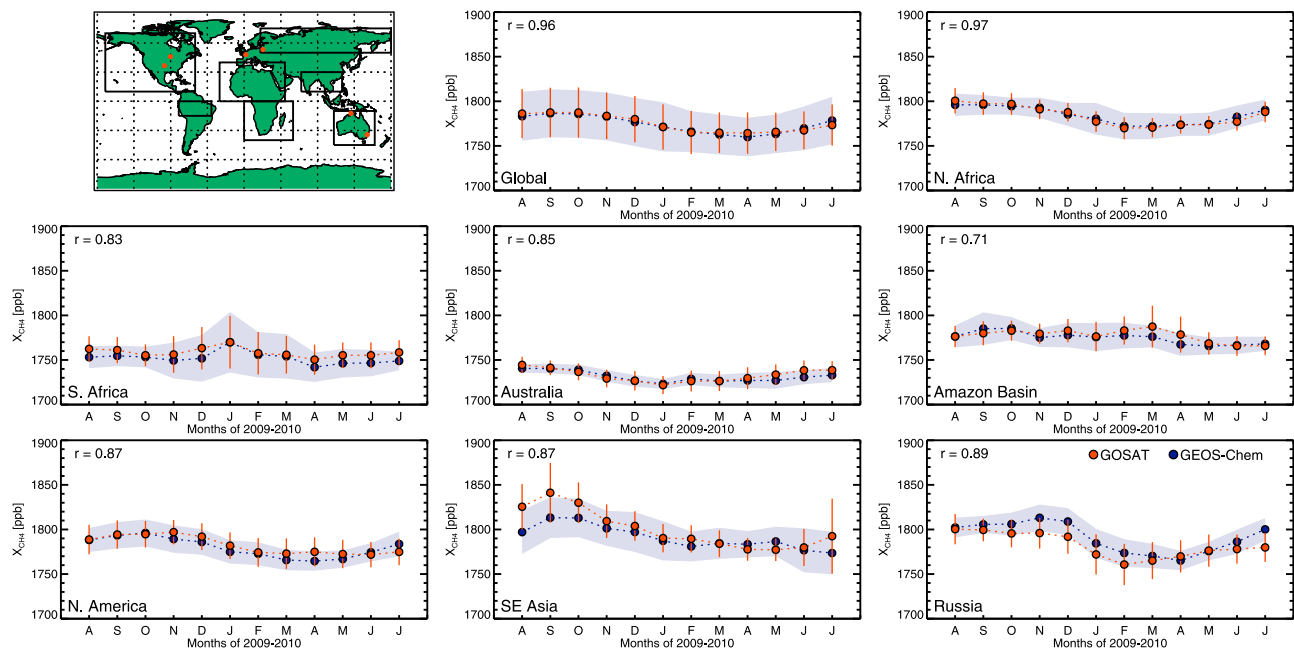
[30] The strongest differences are observed for our Russia region, which extends from Finland, across Russia to eastern Siberia (incorporating boreal wetlands and biomass burning) and the Southeast Asia region, which includes emissions from rice cultivation over India and biomass burning. These CH<sub>4</sub> sources have a large seasonal pattern and we observe differences in the magnitude of these seasonal variations between model and observations of up to 20 ppb for the Russia region and up to 30 ppb for Southeast Asia, which reflects the high level of uncertainty of the emission inventories for this region. Despite these differences in the magnitude of the seasonal signals between measurement and model, the model is still able to capture 70% of the observed variability.

## 6. Concluding Remarks

[31] We present global CH<sub>4</sub> retrievals from GOSAT for August 2009 to July 2010 which we evaluate against correlative measurements from ground-based TCCON sites and calculations from the global 3-D GEOS-Chem chemistry transport model.

[32] From comparisons to TCCON observations we have inferred a single sounding precision for our CH<sub>4</sub> retrievals of 0.4 – 0.8% with estimated biases between  $-17$  ppb and 2 ppb (0.1 to  $-0.9\%$ ) which is well within the mission targets of the GOSAT mission. We expect that the biases in the





**Figure 3.** Time-series of the GOSAT and GEOS-Chem  $X_{CH_4}$  between August 2009 and July 2010 globally and for the 7 regions outlined on the map. The error bars and shaded area represent the standard deviation of the GOSAT and GEOS-Chem data respectively. The location of TCCON sites used in section 4 are indicated by red dots.

retrieval can be reduced in the future by improvements in spectroscopy and instrument calibration. Potential biases in our retrieval might also be introduced by the use of a model CO<sub>2</sub> field as is required by the adopted proxy retrieval approach. Future work will have to focus on further improvements and continued validation of the  $X_{CH_4}$  retrieval, but we expect that the performance of our current retrieval will be sufficient to provide information on the spatio-temporal distribution of CH<sub>4</sub> sources which can lead to improved estimates of CH<sub>4</sub> surface fluxes.

[33] We find a high level of consistency between our GOSAT  $X_{CH_4}$  and large-scale features known to be well-represented by the model, such as the north-south gradient ( $0.93 \leq r \leq 0.99$ ) and seasonal cycle over regions without significant sources (such as northern Africa where  $r = 0.97$ ). While the model typically reproduces >70% of the variability observed with GOSAT, there are however clear differences in the magnitude of the seasonal patterns over known source regions (such as the Amazon ( $r = 0.71$ )) where the emission inventories are highly uncertain. These can be investigated further using an inverse model to quantify the corresponding flux estimates.

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